### Towards 2D overland flow simulations

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# Problem context



### Preventing overland flow and erosion

From upstream...



(Photos : Yves Le Bissonnais, INRA)

#### ...to downstream.





### Downstream zones modifications (watersheds)



- Where is the water coming from ?
- Where is it flowing?

Use of physical models is required to :

- simulate flow (volumes and location)
- suggest changes (grass strip).
- $\rightsquigarrow$  to carry out improvements

# Shallow Water (Saint-Venant) system



Data : topography z, rain P, infiltration IUnknowns : velocities u, v, water height h

$$\begin{cases} \partial_t \mathbf{h} + \partial_x (hu) + \partial_y (hv) = P - I \\ \partial_t (hu) + \partial_x (hu^2 + gh^2/2) + \partial_y (huv) = gh(-\partial_x z - S_{f_x}) \\ \partial_t (hv) + \partial_x (huv) + \partial_y (hv^2 + gh^2/2) = gh(-\partial_y z - S_{f_y}) \end{cases}$$

# Strategy

- Properties of the 1D Shallow Water system
- Choice of the method depending on the properties
- Validation : analytical solutions and laboratory experiment
- Application : field data

## 1D Shallow Water system



$$\partial_t (hu) + \partial_x (hu) - F - F$$

$$\partial_t (hu) + \partial_x (hu^2 + gh^2/2) = gh(-\partial_x z - S_f)$$
(1)

# System properties (I) : Hyperbolicity

Setting q = hu

$$U = \begin{pmatrix} h \\ q \end{pmatrix}, F(U) = \begin{pmatrix} q \\ q^2/h + gh^2/2 \end{pmatrix}, B = \begin{pmatrix} P - I \\ gh(-\partial_x z - S_f) \end{pmatrix},$$

compact form

$$\partial_t \mathbf{U} + \partial_x F(\mathbf{U}) = \partial_t \mathbf{U} + F'(\mathbf{U})\partial_x \mathbf{U} = \mathbf{B},$$

Hyperbolicity if h > 0:

$$\lambda_{-}(U) = u - \sqrt{gh}, \quad \lambda_{+}(U) = u + \sqrt{gh}$$

Saint-Venant	gaz dynamic
Froude number $Fr = rac{ u }{c}$ $c = \sqrt{gh}$ free surface waves celerity	Mach number $\frac{ u }{c}$ $c = \sqrt{p'(\rho)}$ sound speed <sup>1</sup>
subcritical $Fr < 1$	subsonic
supercritical $Fr > 1$	supersonic

1.  $p(\rho) = \rho RT$  perfect gaz

# System properties (II) : Conservation laws

Integral of equation (1) in x

$$\partial_t \mathbf{h} + \partial_x \mathbf{q} = \mathbf{P} - \mathbf{I},$$

gives

$$\frac{d}{dt}\int_a^b h(t,x)\,dx = q(t,a) - q(t,b) + \int_a^b P(t,x) - I(t,x)dx,$$

Mass conservation of water.

Second equation : momentum equation

# System properties (III) : Steady states



# System properties (III) : Steady state

$$\begin{cases} \partial_t h + \partial_x(hu) = P - I \\ \partial_t(hu) + \partial_x(hu^2 + gh^2/2) = gh(-\partial_x z - S_f) \\ \partial_t h = \partial_t u = \partial_t q = 0 \\ \begin{cases} \partial_x hu = P - I \\ \partial_x(hu^2 + gh^2/2) = gh(-\partial_x z - S_f) \end{cases}$$
(2)

# System properties (III) : Steady states

Lac at rest equilibrium

$$u = 0$$
$$g(h + z) = Cst$$

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# Numerical method (I)



$$U_i^{n+1} = U_i^n - \frac{\Delta t}{\Delta x} \big[ F_{i+1/2}^n - F_{i-1/2}^n \big],$$

with the interface flux approximation

$$F_{i+1/2}^n = \mathcal{F}(U_i^n, U_{i+1}^n) \sim \frac{1}{\Delta t} \int_{t^n}^{t^{n+1}} F(U(t, x_{i+1/2})) dt$$

Numerical method (I)

▶ For each choice of  $\mathcal{F}(U_G, U_D)$  we have a different finite volume scheme :

HLL, kinetic, Rusanov, VFRoe-ncv, suliciu, ...

- second Order
  - in space : MUSCL, ENO, modified ENO
  - ▶ in time : Heun

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- second Order
  - in space : MUSCL, ENO, modified ENO
  - in time : Heun
- Coupling with the source term (topography ∂<sub>x</sub>z) Necessity : compatibility with steady states

Steady states (II)

$$\begin{cases} \partial_t h + \partial_x (hu) = 0\\ \partial_t (hu) + \partial_x (hu^2 + gh^2/2) = -gh\partial_x z\\ \partial_t h = \partial_t u = \partial_t q = 0\\ \begin{cases} hu = Cst\\ u^2/2 + g(h+z) = Cst \end{cases} \end{cases}$$
(3)

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We consider

$$\begin{cases} u = Cst \\ g(h+z) = Cst \end{cases}$$

# Hydrostatic reconstruction (II) [Audusse et al., 2004]

We define

$$z^* = max(z_G, z_D)$$

and

$$\begin{cases} U_G^* = (h_G^*, h_G^* u_G), \ U_D^* = (h_D^*, h_D^* u_D) \\ h_G^* = max(h_G + z_G - z^*, 0) \\ h_D^* = max(h_D + z_D - z^*, 0) \end{cases}$$

Thus, we have

$$\begin{cases} \mathcal{F}_{G}(U_{G}, U_{D}, \Delta Z) = \mathcal{F}(U_{G}^{*}, U_{D}^{*}) + \begin{pmatrix} 0 \\ g(h_{G}^{2} - (h_{G}^{*})^{2})/2 \end{pmatrix} \\ \mathcal{F}_{D}(U_{G}, U_{D}, \Delta Z) = \mathcal{F}(U_{G}^{*}, U_{D}^{*}) + \begin{pmatrix} 0 \\ g(h_{D}^{2} - (h_{D}^{*})^{2})/2 \end{pmatrix} \end{cases},$$

where  $\mathcal{F}(U_G, U_D)$  is the numerical flux.

### Friction treatment

Shallow Water system with friction f

$$\begin{cases} \partial_t \mathbf{h} + \partial_x(\mathbf{h}\mathbf{u}) = 0, \\ \partial_t(\mathbf{h}\mathbf{u}) + \partial_x(\mathbf{h}\mathbf{u}^2 + \mathbf{g}\mathbf{h}^2/2) + \mathbf{h}\partial_x z = -\mathbf{h}f, \end{cases}$$
(4)

f = f(h, u) friction force (on the bottom)

Several friction laws possible

- Manning :  $f = n^2 \frac{u|u|}{h^{4/3}}$
- Darcy-Weisbach :  $f = F \frac{u|u|}{8gh}$

Apparent topography [Bouchut, 2004]

We consider :  $z_{app} = z + b^n$ 

with  $\partial_x b^n = S_f^n$ 

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Semi-implicit [Bristeau and Coussin, 2001]

$$q_{i}^{n+1} + F \frac{|q_{i}^{n}|q_{i}^{n+1}}{8h_{i}^{n}h_{i}^{n+1}} \Delta t = q_{i}^{n} + \frac{\Delta t}{\Delta x} \left(F_{i+1/2} - F_{i-1/2}\right)$$

with  $q_i^{n+1*}$  for the right part, we have

$$q_i^{n+1} = rac{q_i^{n+1*}}{1+\Delta t rac{F|u_i^n|}{8h_i^{n+1}}}$$

New test cases :

- Saint-Venant/shallow water :
  - data z
  - unknowns h et u (and so q)

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- Saint-Venant/shallow water :
  - data z
  - unknowns h et u (and so q)
- test cases
  - data h and q (and so u)
  - unknown z
- Several possibilities
  - several friction laws
  - diffusion source term [Delestre and Marche, 2010]
  - rain source term



Apparent topography (subcritical-subcritical)



Semi-implicit (subcritical-subcritical)



Semi-implicit (subcritical-supercritical)



Semi-implicit (supercritical-subcritical)



### Summary of the chosen numerical method

- Numerical flux : HLL
- Second order scheme : MUSCL
- Friction : semi-implicit treatment
- Shallow Water system with rain P

$$\begin{cases} \partial_t h + \partial_x (hu) = P \\ \partial_t (hu) + \partial_x (hu^2 + gh^2/2) + h\partial_x z = -hf \end{cases}$$
(5)

time splitting/explicit treatment

# Validation on experiments - INRA rain simulator



## Settings of the experiment



 $0 \leq t \leq 250 \mathrm{s}$ 

$$R(x, t) = \begin{cases} 50 \text{ mm/h if } (x, t) \in [0, L] \times [5, 125] \\ 0 \text{ else} \end{cases}$$

### Analytical solutions and simulations



## Water height and velocity at equilibrium



# What about reality?



## "Calibration"



# A simulation result (Manning)



# Parcels in Niger ([Esteves et al., 2000], IRD)



# Darcy-Weisbach



# Manning



Velocity measures by SVG [Planchon et al., 2005] Number of cells : 40  $\times$  100 (4 m  $\times$  10 m)









Hyetogramme

Number of cells : 160  $\times$  200



Homogeneous coefficients



Homogeneous coefficients



Velocities



Heterogeneous coefficients



#### Velocities



# FullSWOF

- object and inheritance
- variables encapsulation
- vector class (2d)
- objects "distributor"
- fixed CFL and fixed  $\Delta t$
- Doxygen documentation
- Free open source software

### Malpasset dam break simulation (Cordier et al., CEMRACS 2012)



### Malpasset dam break simulation





Maximum water elevation

# Thank You!

HLL flux

$$\mathcal{F}(U_G, U_D) = \begin{cases} F(U_G) \text{ if } 0 < c_1 \\ F(U_D) \text{ if } c_2 < 0 \\ \frac{c_2 F(U_G) - c_1 F(U_D)}{c_2 - c_1} + \frac{c_1 c_2 (U_D - U_G)}{c_2 - c_1} \text{ else} \end{cases},$$

with two parameters

 $c_1 < c_2$ .

For  $c_1$  and  $c_2$ , we take

$$c_1 = \inf_{U = U_G, U_D} (\inf_{j \in \{1,2\}} \lambda_j(U)) \text{ and } c_2 = \sup_{U = U_G, U_D} (\sup_{i \in \{1,2\}} \lambda_i(U)).$$

with  $\lambda_1(U) = u - \sqrt{gh}$  and  $\lambda_2(U) = u + \sqrt{gh}$ . For each  $e^{-\pi i t}$ 

### Audusse, E., Bouchut, F., Bristeau, M.-O., Klein, R., and Perthame, B. (2004).

A fast and stable well-balanced scheme with hydrostatic reconstruction for shallow water flows.

SIAM J. Sci. Comput., 25(6) :2050-2065.



Bouchut, F. (2004).

Nonlinear stability of finite volume methods for hyperbolic conservation laws, and well-balanced schemes for sources, volume 2/2004.

Birkhäuser Basel.

```
Bristeau, M.-O. and Coussin, B. (2001).
```

Boundary conditions for the shallow water equations solved by kinetic schemes.

Technical Report 4282, INRIA.



Delestre, O. and Marche, F. (2010).

A numerical scheme for a viscous shallow water model with friction. J. Sci. Comput., DOI 10.1007/s10915-010-9393-y.



Esteves, M., Faucher, X., Galle, S., and Vauclin, M. (2000). Overland flow and infiltration modelling for small plots during unsteady rain : numerical results versus observed values. Journal of Hydrology, 228 :265-282.

 Planchon, O., Silvera, N., Gimenez, R., Favis-Mortlock, D., Wainwright, J., Le Bissonnais, Y., and Govers, G. (2005).
 An automated salt-tracing gauge for flow-velocity measurement. Earth Surface Processes and Landforms, 30(7) :833–844.



Tatard, L., Planchon, O., Wainwright, J., Nord, G., Favis-Mortlock, D., Silvera, N., Ribolzi, O., Esteves, M., and Huang, C.-h. (2008). Measurement and modelling of high-resolution flow-velocity data under simulated rainfall on a low-slope sandy soil.

Journal of Hydrology, 348(1-2) :1–12.