

# Sediment Transport , Numerical Modeling and Reservoir Management

some Concepts and Applications

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# Overview of the presentation

- ▶ What is Sediment Transport ?
- ▶ Sediment Transport Modeling
- ▶ How to Deal with Sediments in Reservoirs ?
- ▶ Example : Modeling Reservoir Emptying
- ▶ Examples of Current Research : 2011 & 2013 Cemracs Projects



# What kind of sediments do we find in rivers ?

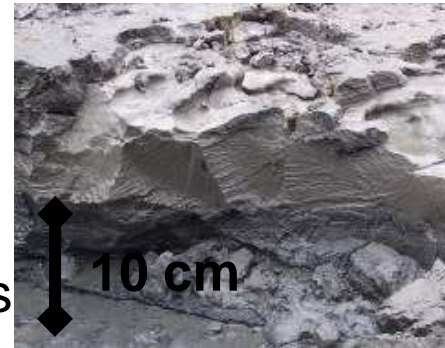
## ▶ Granular Material

- gravel, sand
- non cohesive material
- grain size  $> 40\mu\text{m}$

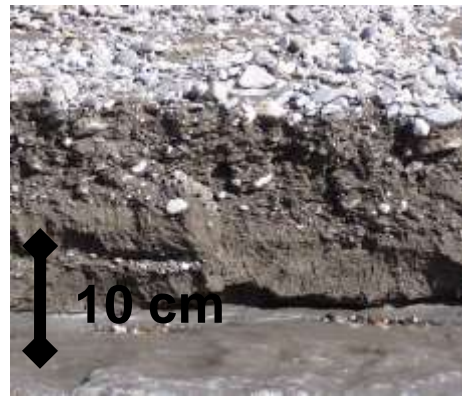


## ▶ Cohesive Material

- Silt , clay
- Grain size  $< 40\mu\text{m}$
- Strong Interactions between particles  
→ Cohesion and flocculation



## ▶ Mixing of gravels and silt :

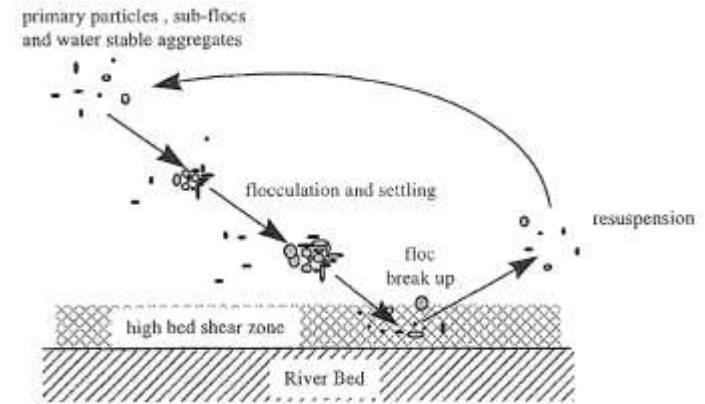


# Cohesive / non cohesive → different physical properties

► Flocculation : cohesive sediments may form aggregates

► Consolidation of cohesive sediments

► Bank stability : different kind of stabilities



Cohesive



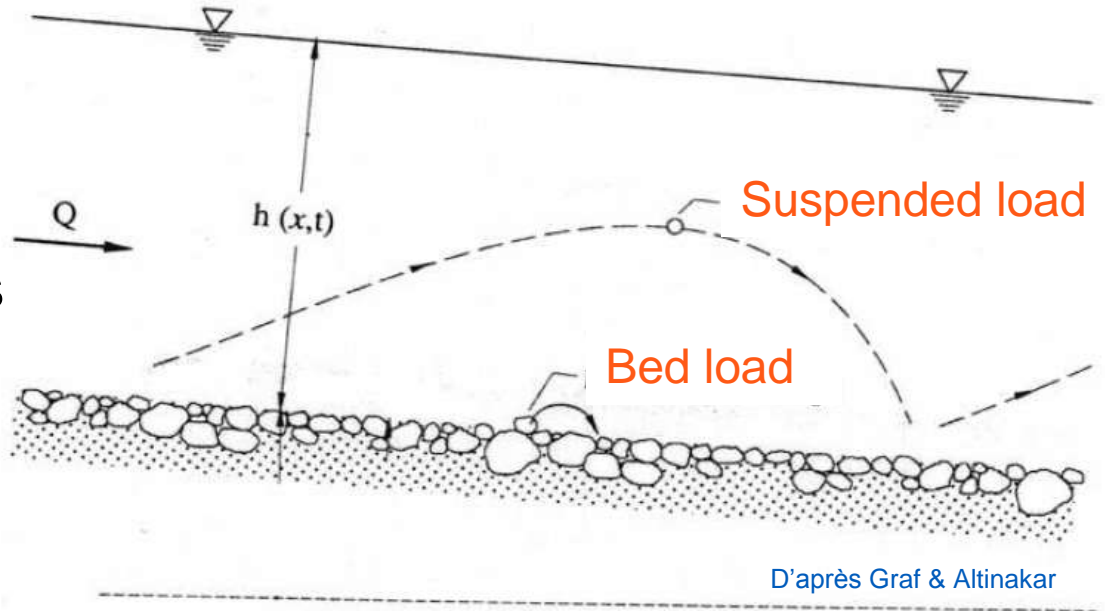
Non cohesive

# Transport of Sediments in Rivers

## Sand and gravel

### → Bed load Transport

saltating and rolling  
near the bed of sediments



## Fine sediments

### → Suspended transport

mixing of sediments in the water

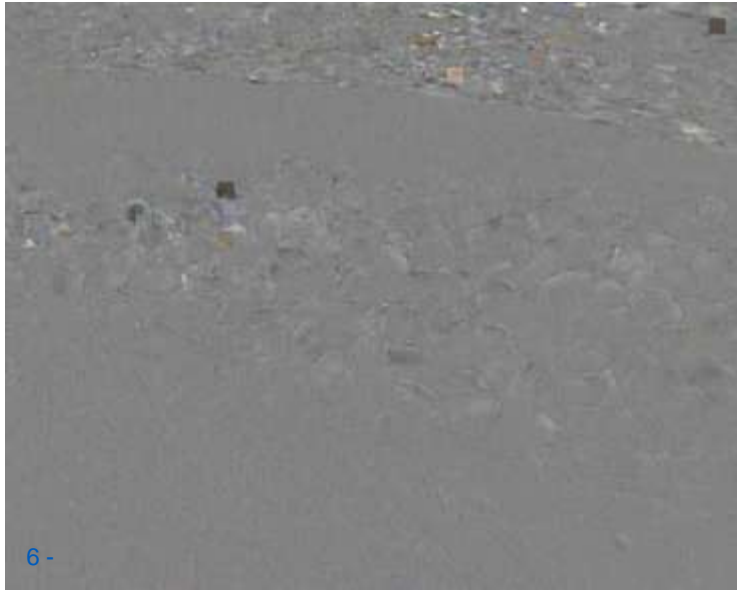
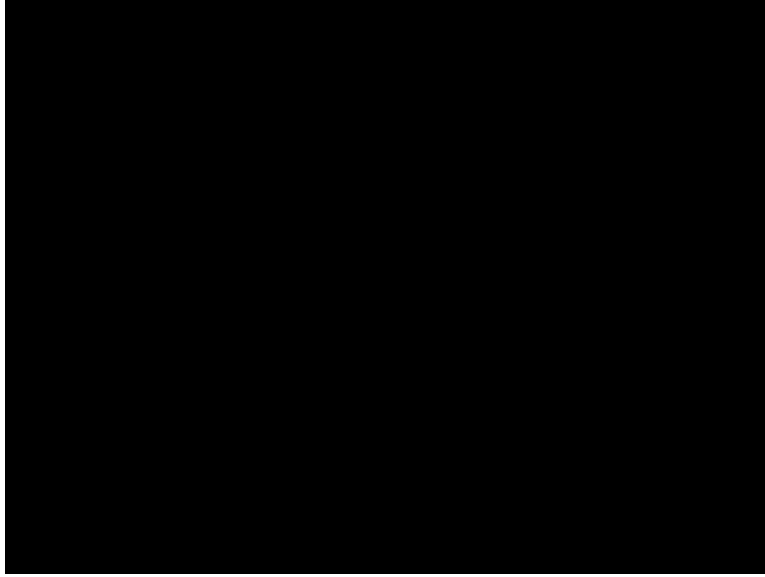
**Advection dispersion equation**



# Transport of Sediments in Rivers

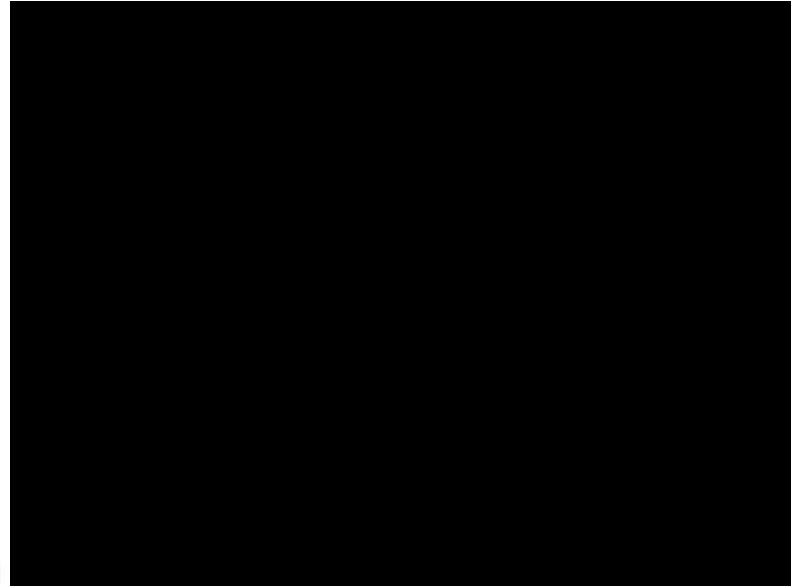
**Sand and gravel**

→ **Bed load Transport**

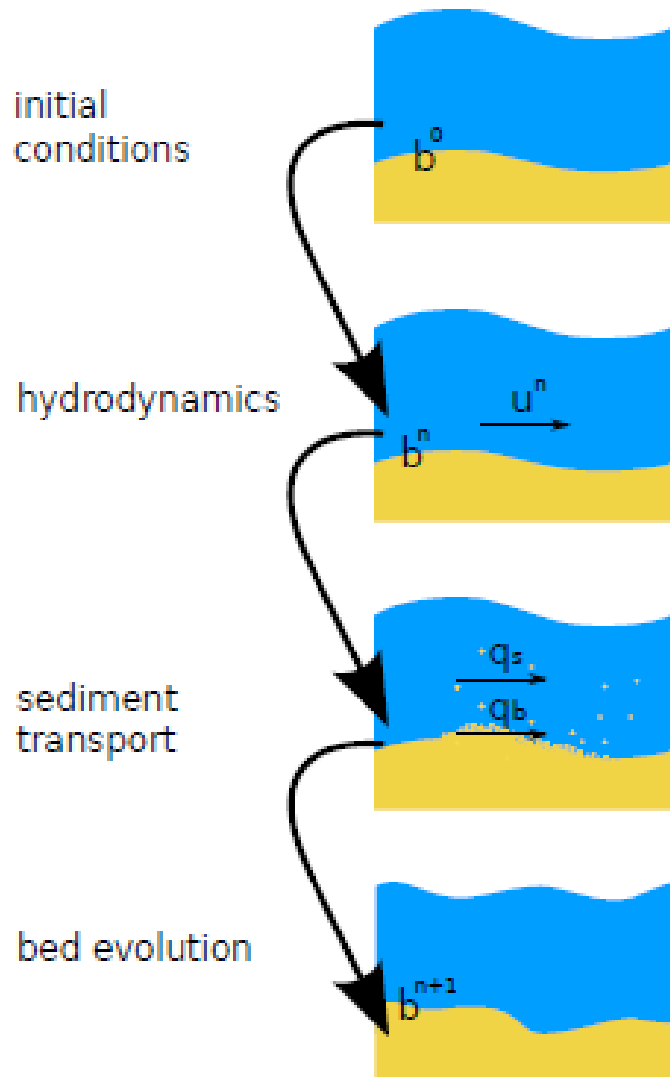


**Fine sediments**

→ **Suspended transport**



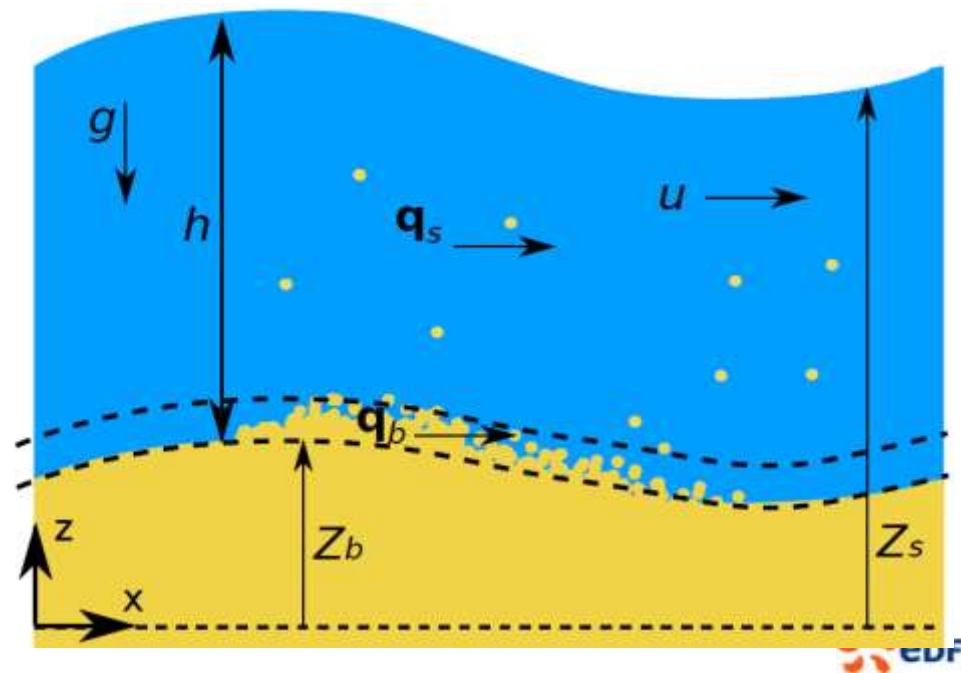
# Sediment Transport Modeling : Processes



→ Need of

1. a set of equations for Hydrodynamics
2. a set of equations for Sediment Transport and Bed Evolution

1 & 2 could be splitted



# Sediment transport modeling : 3D/2D/1D

► Users have to choose the numerical code depending on the goal of the simulation

► 3D, 2D and 1D simulations are possible

► **Empirical formulae** are used for bed interaction and sediment fluxes

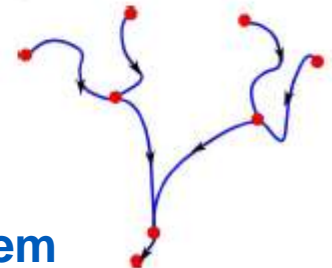
► **Sedi3D**: 3D sediment transport module



► **Sisyphe**: 2D sediment transport module



► **Courlis**: 1D sediment transport module

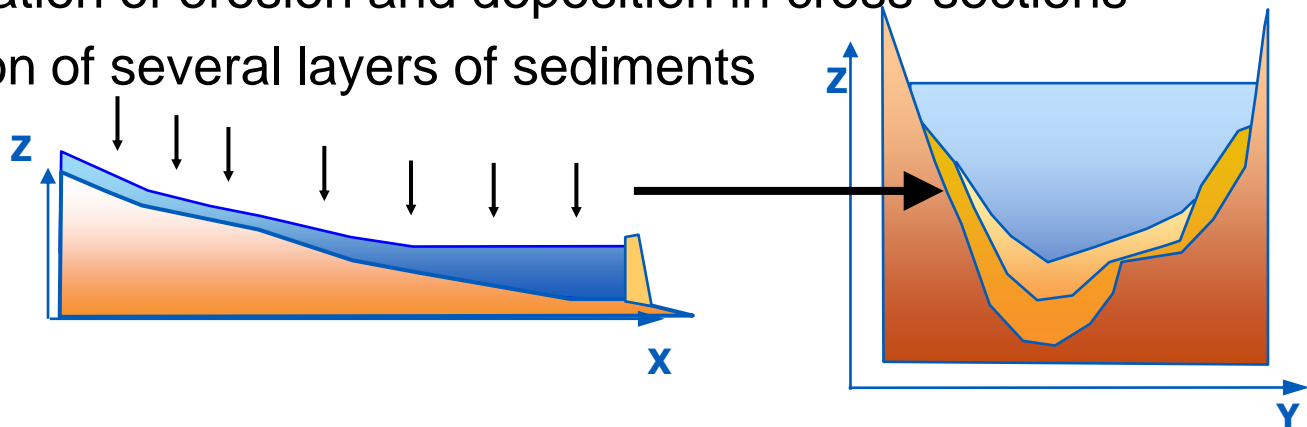


EDF sediment transport tools  
**Open Source Telemac Mascaret System**  
<http://www.opentelemac.org/>



# Example of the 1D sediment transport numerical code

- ▶ COURLIS numerical code (Bertier et al 2002)
- ▶ One dimensional
- ▶ Part of Telemac-Mascaret system (<http://www.opentelemac.org/>)
- ▶ Coupling between 1D shallow water equations (Mascaret, Goutal and Maurel 2002) and sediment component : Splitting approach
- ▶ 2D calculation of erosion and deposition in cross-sections
- ▶ Description of several layers of sediments



# COURLIS (Bertier et al 2002)

## ► Hydrodynamics component : Mascaret

It solves the Shallow Water Equations

Mass continuity equation

$$\left\{ \begin{array}{l} \frac{\partial A}{\partial t} + \frac{\partial Q}{\partial x} = 0 \end{array} \right.$$

Momentum equation

$$\left\{ \begin{array}{l} \frac{\partial Q}{\partial t} + \frac{\partial}{\partial x} \left( \frac{Q^2}{A} \right) = -g.A. \frac{\partial Z}{\partial x} - g.A.J \end{array} \right.$$

avec : A la section mouillée (m<sup>2</sup>)

Q le débit (m<sup>3</sup>/s)

Z la cote de la surface libre (m)

J la pente d'énergie déterminée à partir de la relation de Strickler.

## ► Sediment component for SUSPENSION transport (fine sediments) :

- Advection dispersion equation for sand and silt (independent)

$$\frac{\partial AC}{\partial t} + \frac{\partial QC}{\partial x} = \frac{\partial}{\partial x} \left( kA \frac{\partial C}{\partial x} \right) + E - D \quad \text{Bed interactions}$$

- Partheniades and Krone formulae for erosion and deposition of cohesive sediments

$$E = M \left( \frac{\tau}{\tau_{CE}} - 1 \right) \quad D = w_s C \left( 1 - \frac{\tau}{\tau_{CD}} \right)$$

- Engelund Hansen Formula for sand transport capacity  $q_s = 0.05 \sqrt{\frac{\delta d^3}{g} \frac{K^2 R_h^{1/3} \tau_{eff}}{(\rho_s - \rho)gd}}$   $C_{eq} = \frac{\rho_s q_s}{Q}$

- Erosion and deposition rates  $\begin{cases} \text{if } C_{sand} \geq C_{eq} \text{ deposition} & D = w_s (C_{sand} - C_{eq}) \\ \text{if } C_{sand} \leq C_{eq} \text{ erosion} & E = w_s (C_{eq} - C_{sand}) \end{cases}$

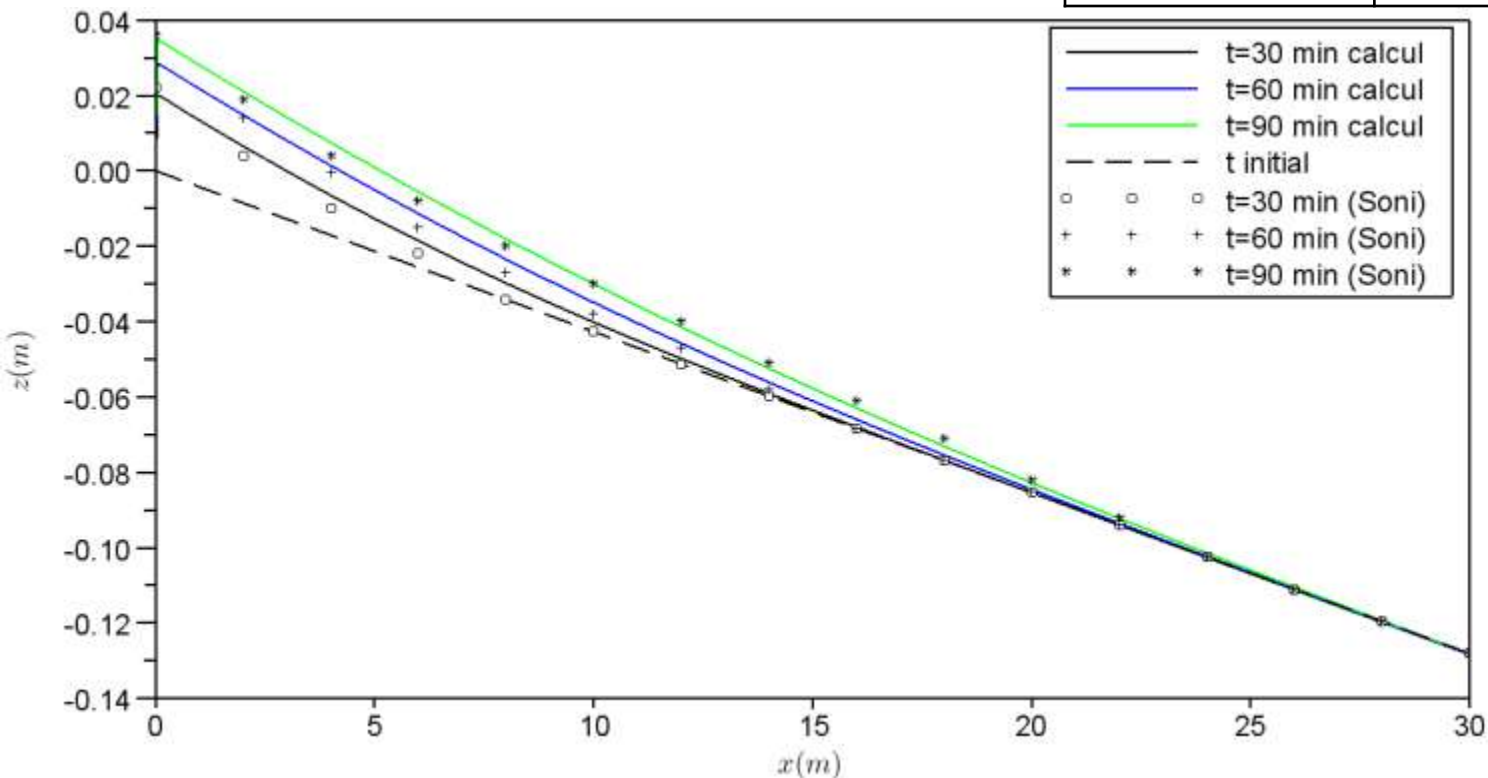
- Bed evolution  $\frac{\partial Zb}{\partial t} = \frac{D}{C_{deposition}} - \frac{E}{C_{layer}}$

# Sand Deposition test case : Soni experiment

Soni J.P. Laboratory study of aggradation in alluvial channels, Journal of Hydrology, (49), 1981

- Mesh size  $\Delta x=25\text{cm}$
- Friction coefficient  $K_s=45\text{ m}^{1/3}\text{s}^{-1}$
- Diffusion coefficient  $K_x=0.025\text{m}^2\text{s}^{-1}$
- Non equilibrium coefficient  $\alpha=0.54$

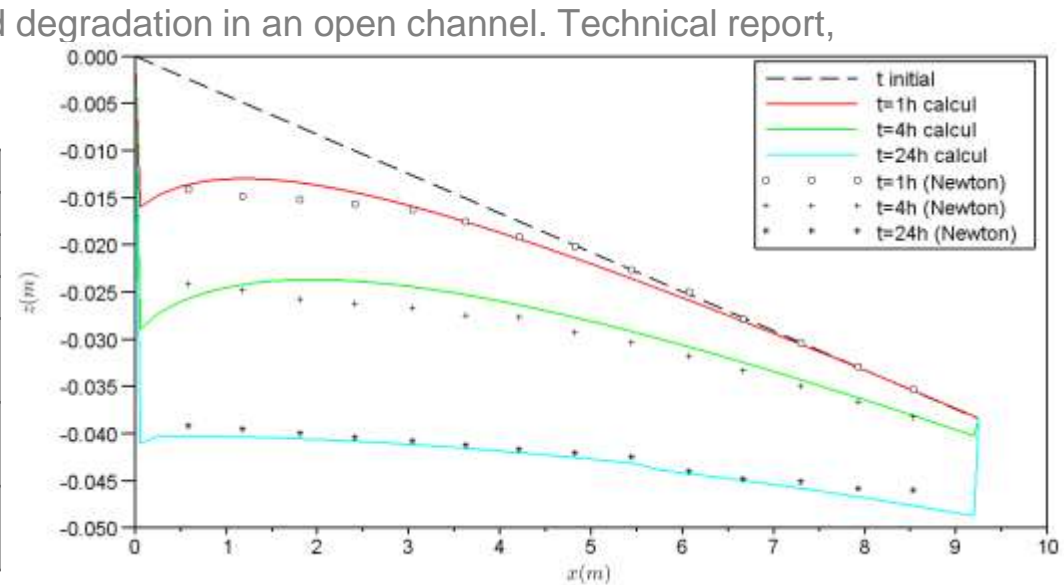
Flume length	L	30	(m)
Flume width	w	20	(cm)
Slope	S	$4.27 \cdot 10^{-3}$	
Discharge		$7.1 \cdot 10^{-3}$	( $\text{m}^3/\text{s}$ )
Downstream water depth	Hd	7.2	(cm)
Upstream concentration	Cu	4.88	(g/l)
Median grain size	$d_{50}$	320	( $\mu\text{m}$ )



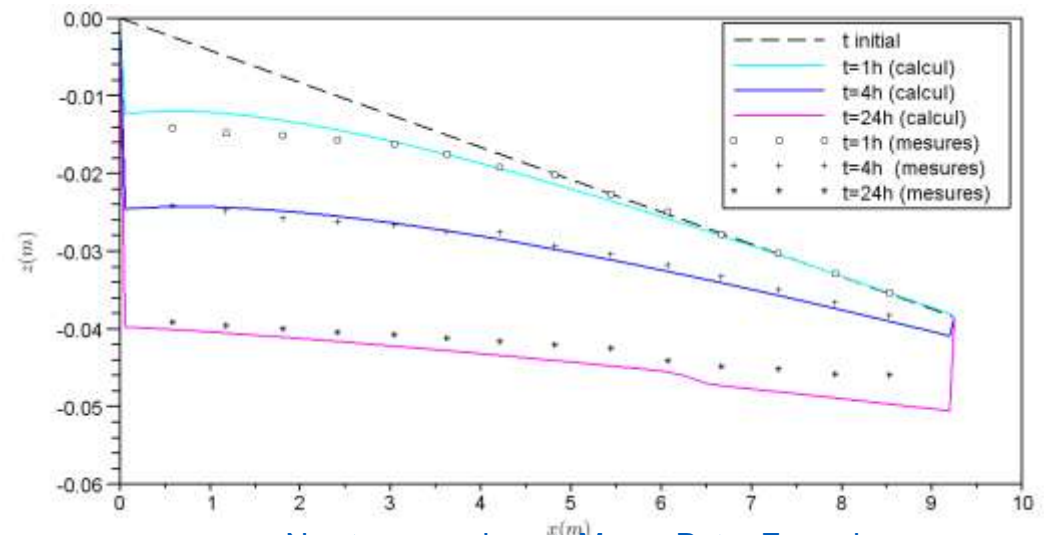
# Sand Erosion test case : Newton Experiment

Newton C.T. An experimental investigation of bed degradation in an open channel. Technical report, Boston Society of Civil Engineers, 1951

Flume length	L	9.14	(m)
Flume width	w	30.48	(cm)
Slope	S	$4.16 \cdot 10^{-3}$	
Discharge		$5.66 \cdot 10^{-3}$	(m <sup>3</sup> /s)
Downstream water depth	Hd	4.1	(cm)
Upstream concentration	Cu	0.88	(g/l)
Median grain size	d <sub>50</sub>	680	(μm)



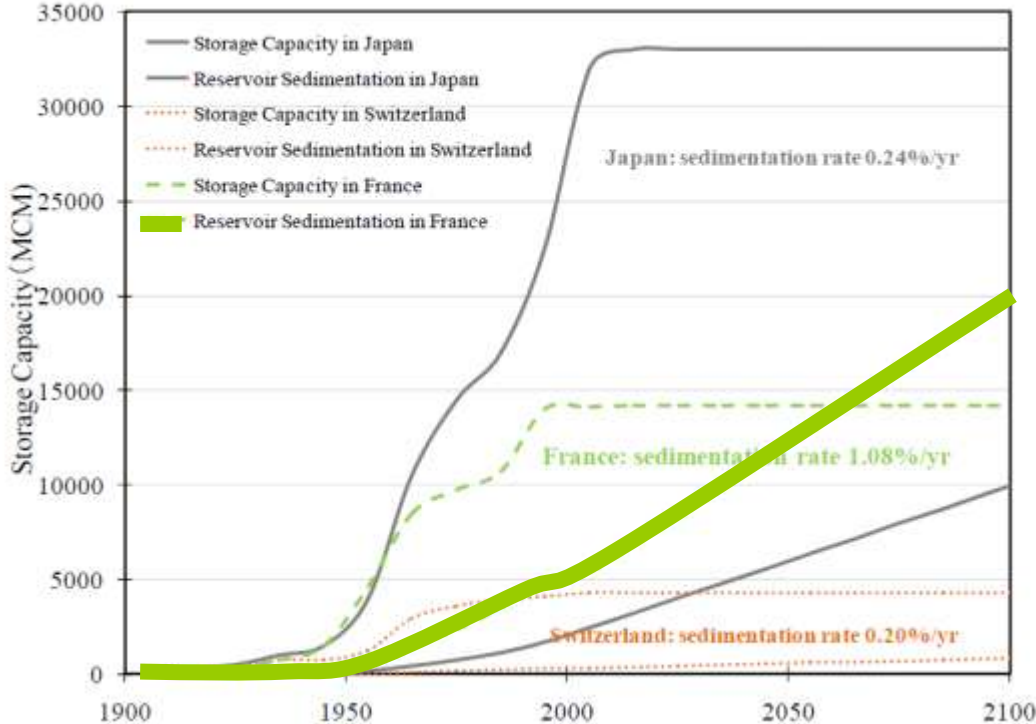
Newton experiment, Enguelund Hansen Formula



Newton experiment, Meyer Peter Formula

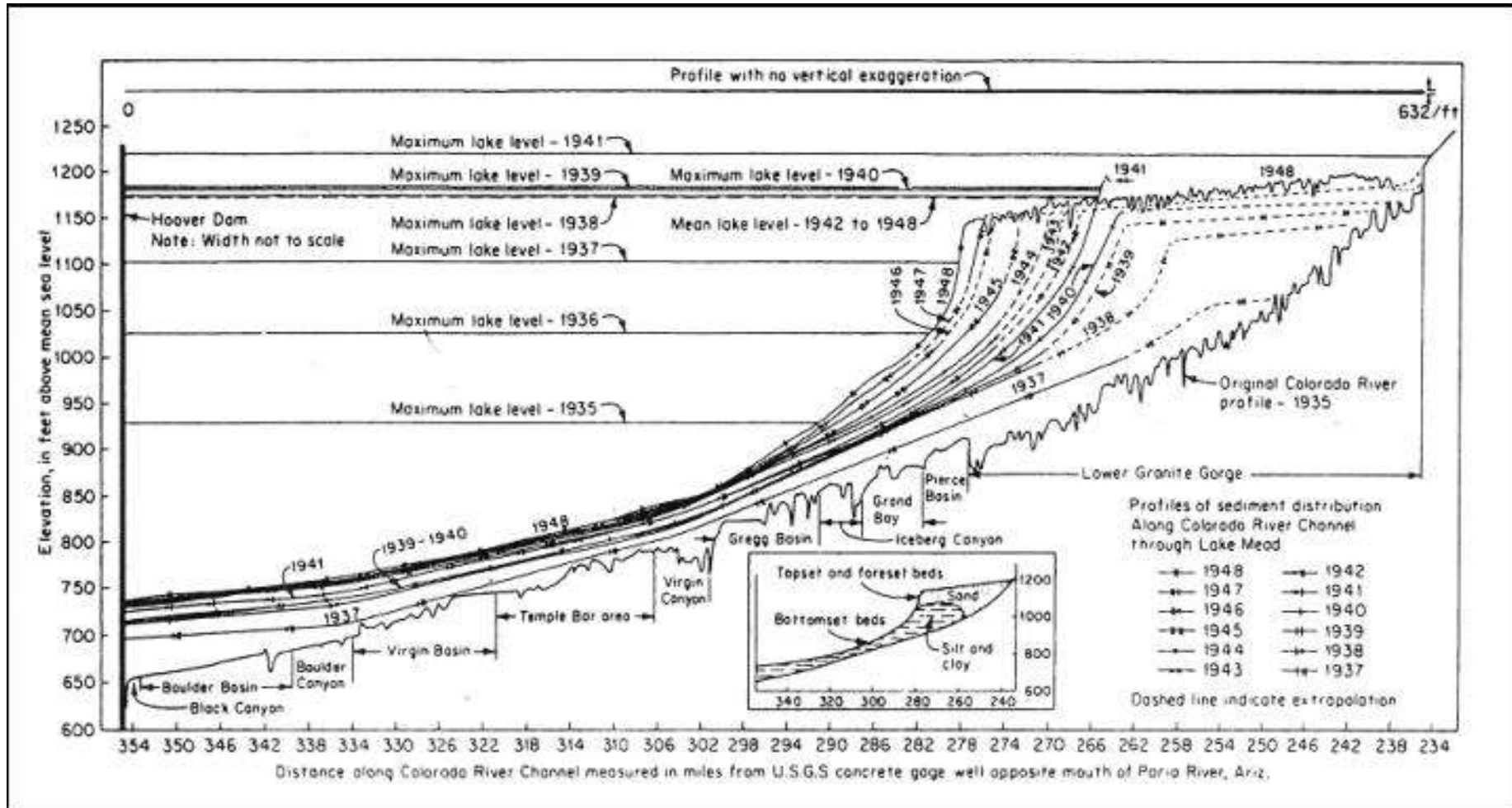
- Mesh size  $\Delta x=25\text{cm}$
- Friction coefficient  $K_s=67 \text{ m}^{1/3}\text{s}^{-1}$
- Diffusion coefficient  $K_x=1\text{m}^2\text{s}^{-1}$

# Sediments in reservoirs



# Sediments in reservoirs

Lac Mead, US ( Smith 1954)



# How to deal with reservoir sedimentation ?

- ▶ EDF manage more than 400 reservoirs
- ▶ **Reservoir emptying** is performed regularly to control the state of dams or to perform works
  - Large quantities of eroded sediments
  - Need to predict downstream impacts (water quality)
- ▶ **Reservoir flushing** is performed to stop reservoir sedimentation
  - Need to know how to manage the flushing
  - Need to forecast downstream transport of sediment
- ▶ **Numerical modeling** a convenient way to deal with these questions



Emptying of Riou reservoir



Swiss Reservoir, picture from T. Bertolcht



# Example :Emptying Tolla Reservoir

Tolla Reservoir (South Corsica) :

- ▶ emptying in order to perform works on the dam
- ▶ mitigate water quality degradation
- ▶ Not so many options : dilution using tributaries, settling tank, **time during the year**, **speed of lowering** and **minimal elevation**
- ▶ **Downstream water intake for drinking water supply for Ajaccio city (53 000 inhabitants)**

Use of numerical modeling

- ▶ to estimate the quantities of eroded sediments
- ▶ To test different scenarios of emptying
- ▶ **Use of a one dimensional model to simulate the downstream concentrations**



Tolla last Emptying1981

# What kind of modeling ?

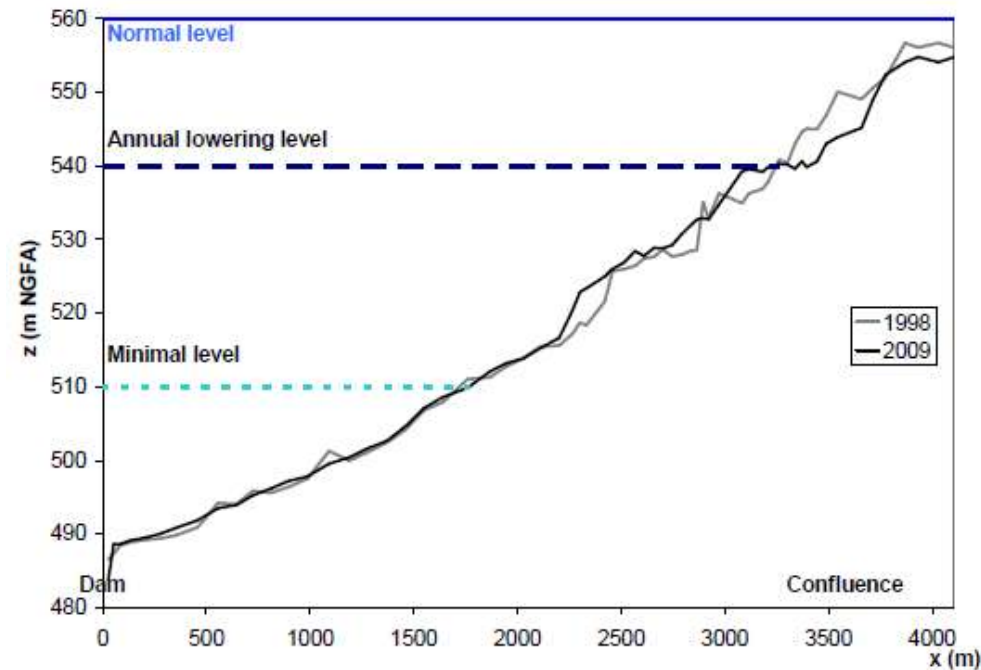
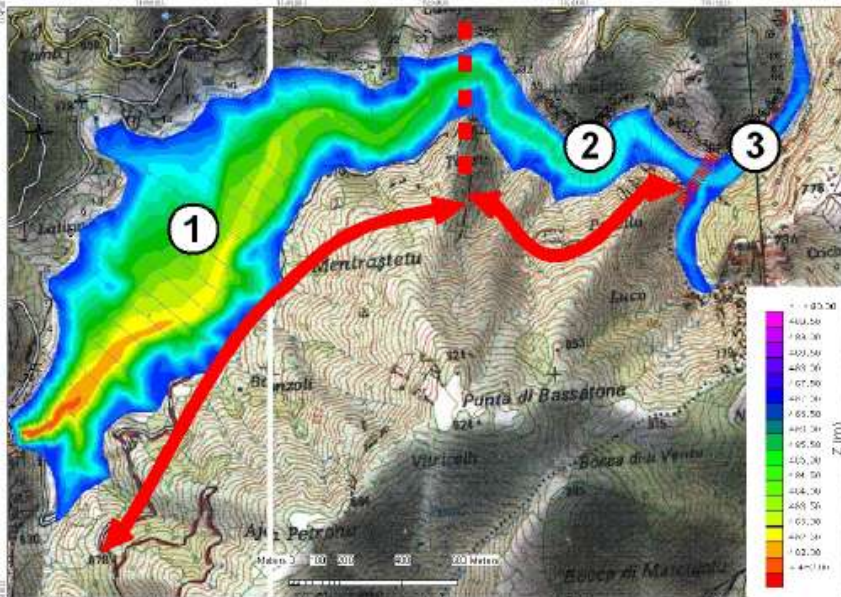
- ▶ **Numerical modeling** a convenient way to deal with reservoir operations and prediction of downstream sediment concentration

One dimensional modeling well suited in many cases

- ▶ Depends on the geometry of the reservoir
- ▶ No need to reproduce in detail flow and sediment transport patterns in the reservoir
- ▶ Very good results for engineering studies on previous cases :
  - St Egreve Reservoir, Valette ICSE 2012,
  - Grangent Reservoir, Bertier River Flow 2012

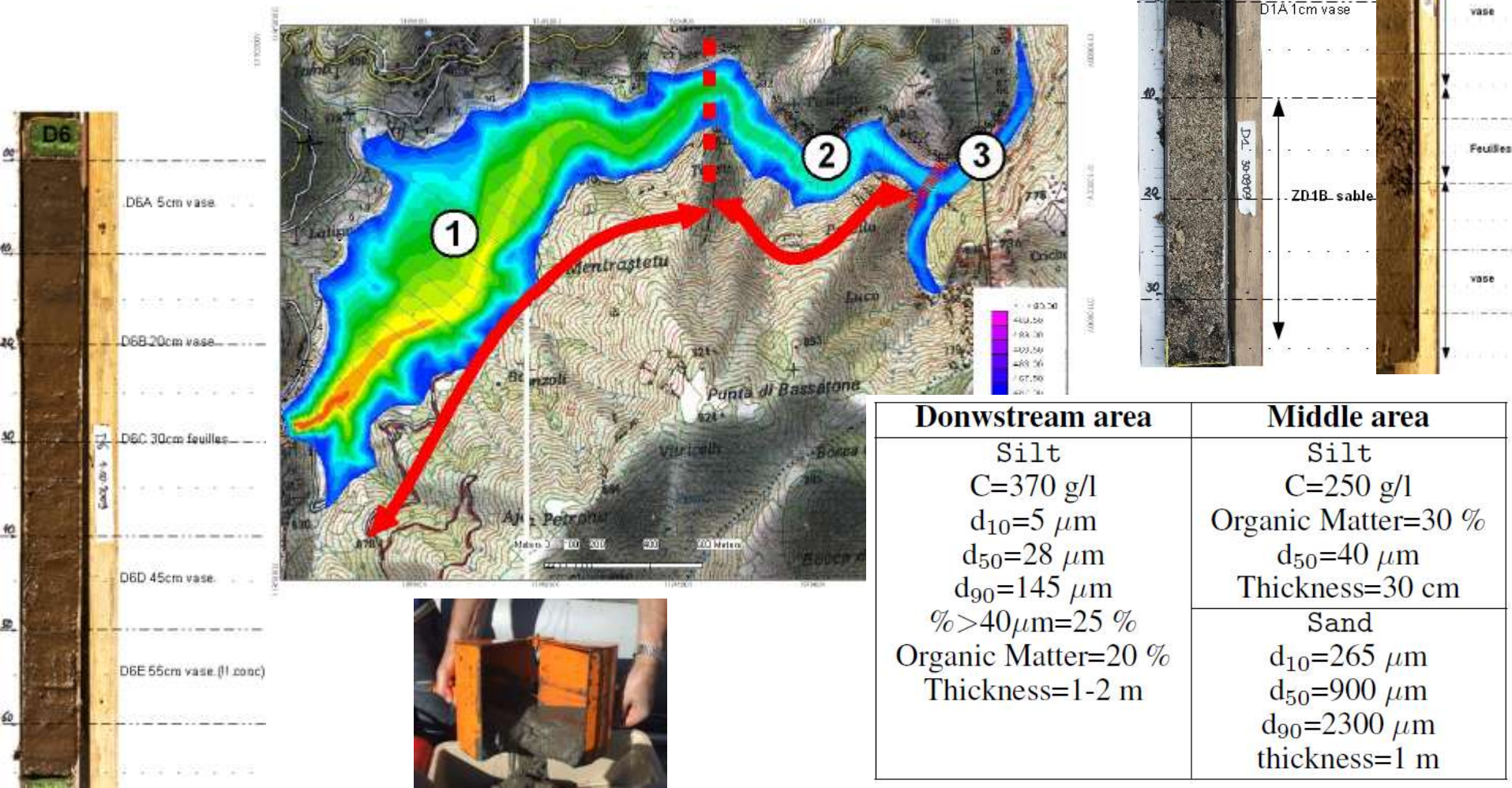
# Sediment and Morphology of Tolla Reservoir

- ▶ 2 bathymetries (1998-2009)
- ▶ Old small dam near the main dam
- ▶ Steep slope x~2200m + upstream confluence



# Sediment properties from sampled cores

- ▶ Silt in the downstream area (1) and sand upstream (2)
- ▶ + leaves
- ▶ Upstream area (3): not modeled



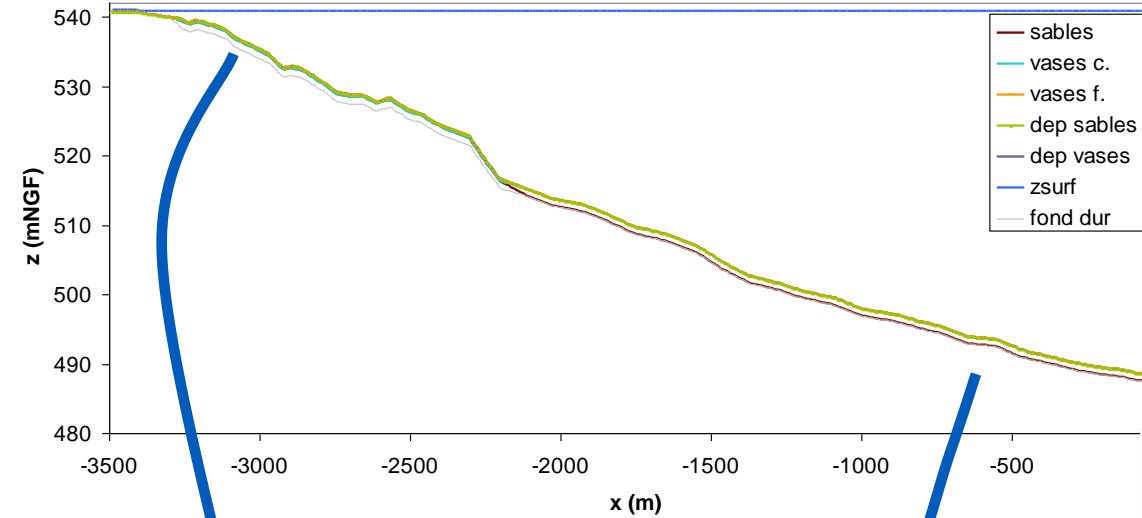
# Description of the reservoir for the model

bed : Bathymetries

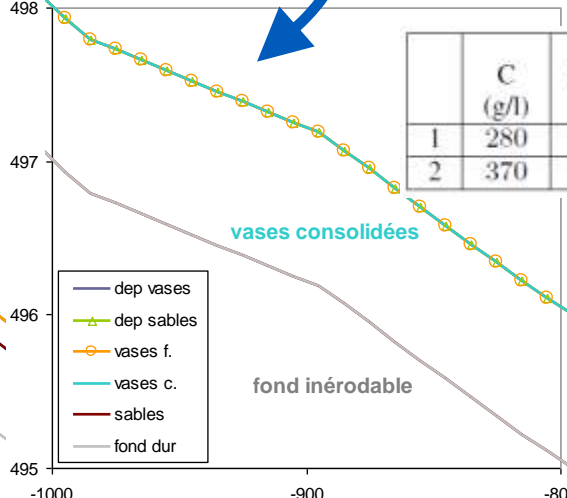
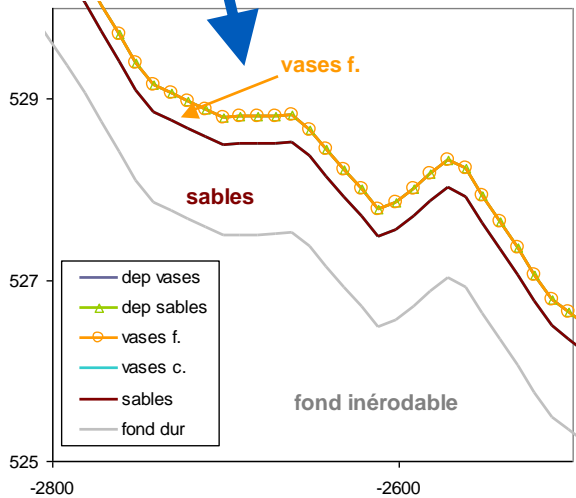
Sediment properties

Lack of calibration data → sensitivity analysis

▶ we choose the worst but physical parameters



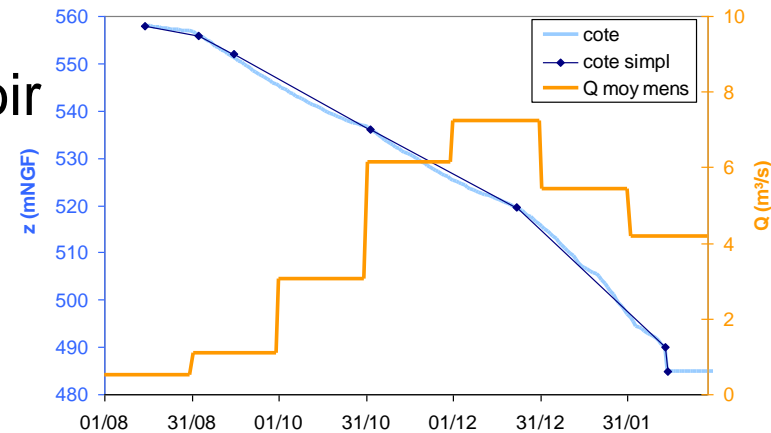
	C (g/l)	Sand parameters			
		sand (%)	$d_{50}$ ( $\mu\text{m}$ )	$w_s$ (m/s)	$K_p$ ( $\text{m}^{1/3}\text{s}^{-1}$ )
sand	1400	90	90	$7 \cdot 10^{-3}$	75 – 95



	C (g/l)	sand (%)	Silt parameters			
			$\tau_{ce}$ (Pa)	M ( $\text{kgs}^{-1}\text{m}^{-2}$ )	$w_s$ (m/s)	$\tau_{cd}$ (Pa)
1	280	0	0.1 – 1	0.01 – 0.02	$2 \cdot 10^{-3}$ – $10^{-2}$	0.01 – 0.1
2	370	19	0.5 – 2	0.01 – 0.02	$2 \cdot 10^{-3}$ – $10^{-2}$	0.01 – 0.1

# Initial conditions and limit conditions

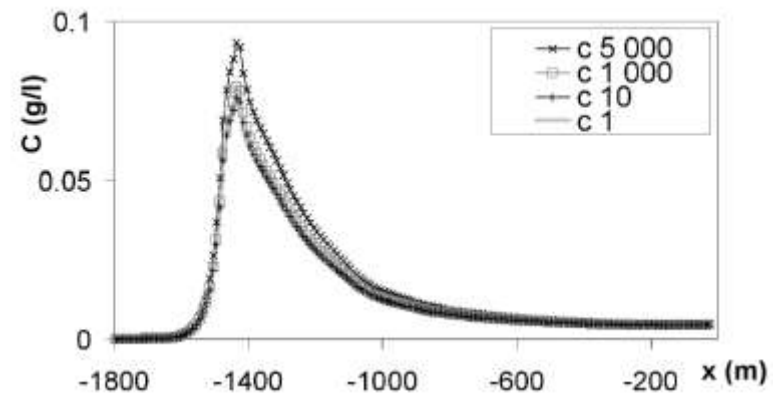
- ▶ Downstream condition : emptying scenario
- ▶ Upstream : incoming discharge
- ▶ Initial : steady state of the full reservoir



# Numerical parameters

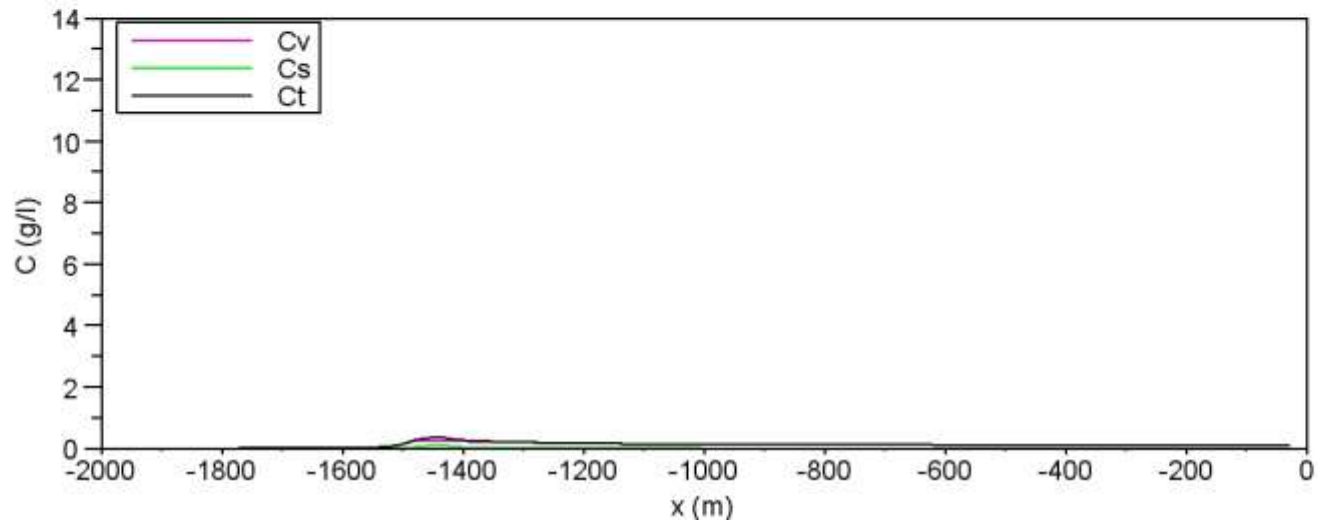
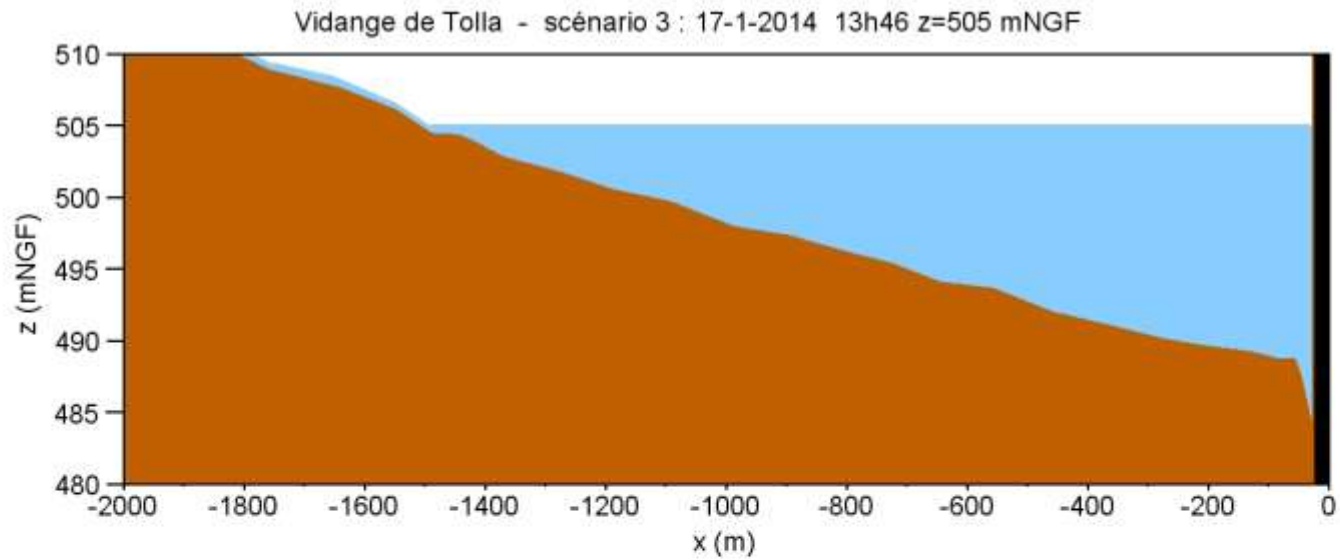
- Vertical and longitudinal meshes
- Numerical schemes (supercritical flows)

- Coupling time step  $\left(u + \sqrt{gh}\right) \frac{\Delta t}{\Delta x} < 1$



- ▶ Chosen to obtain reliable results with smallest calculation times as possible

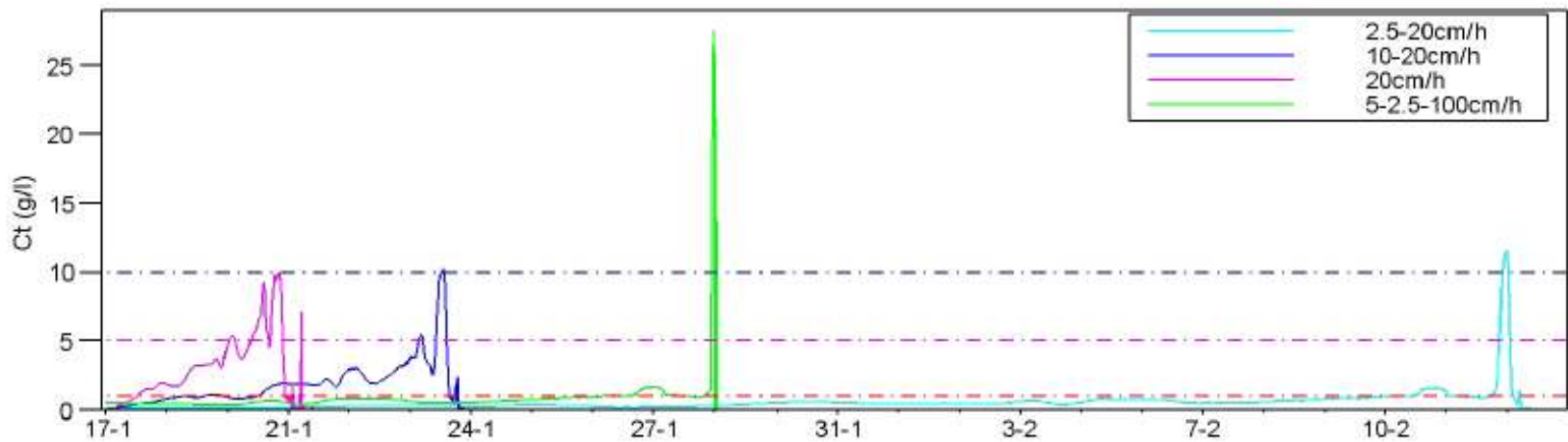
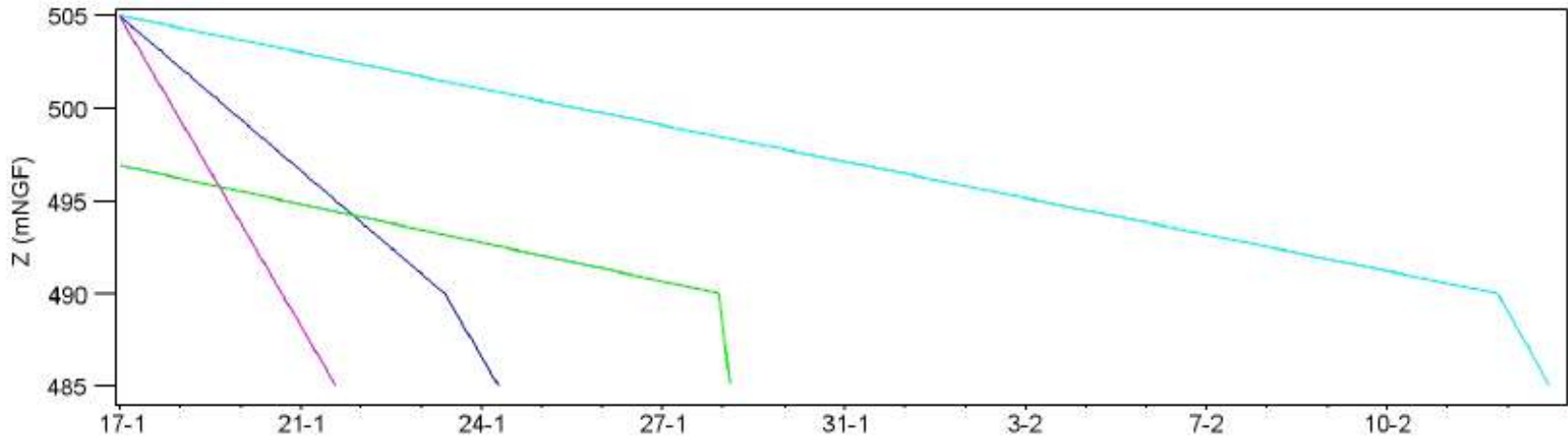
# Results



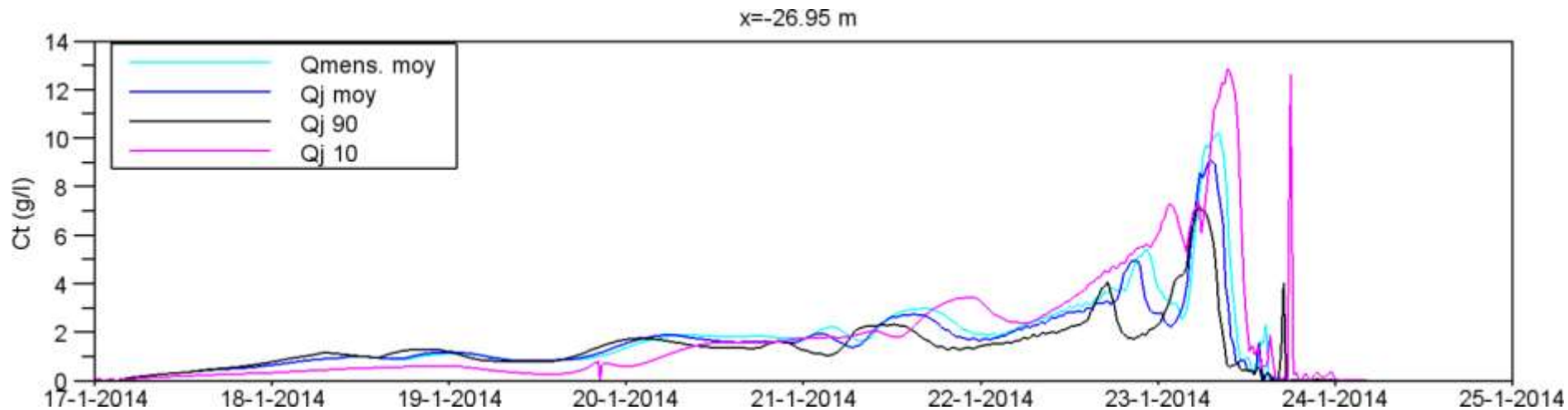
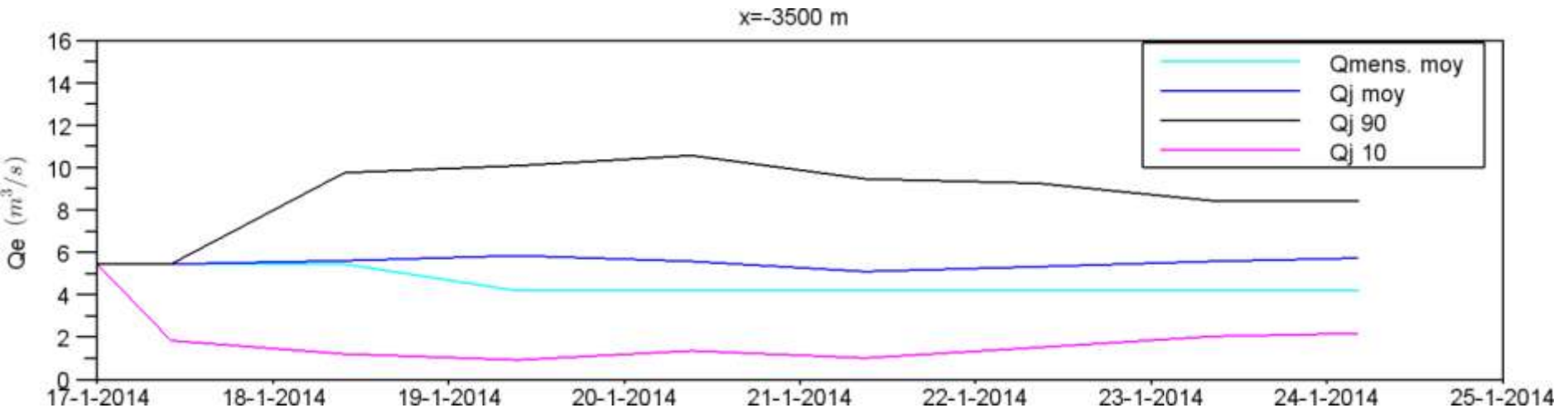




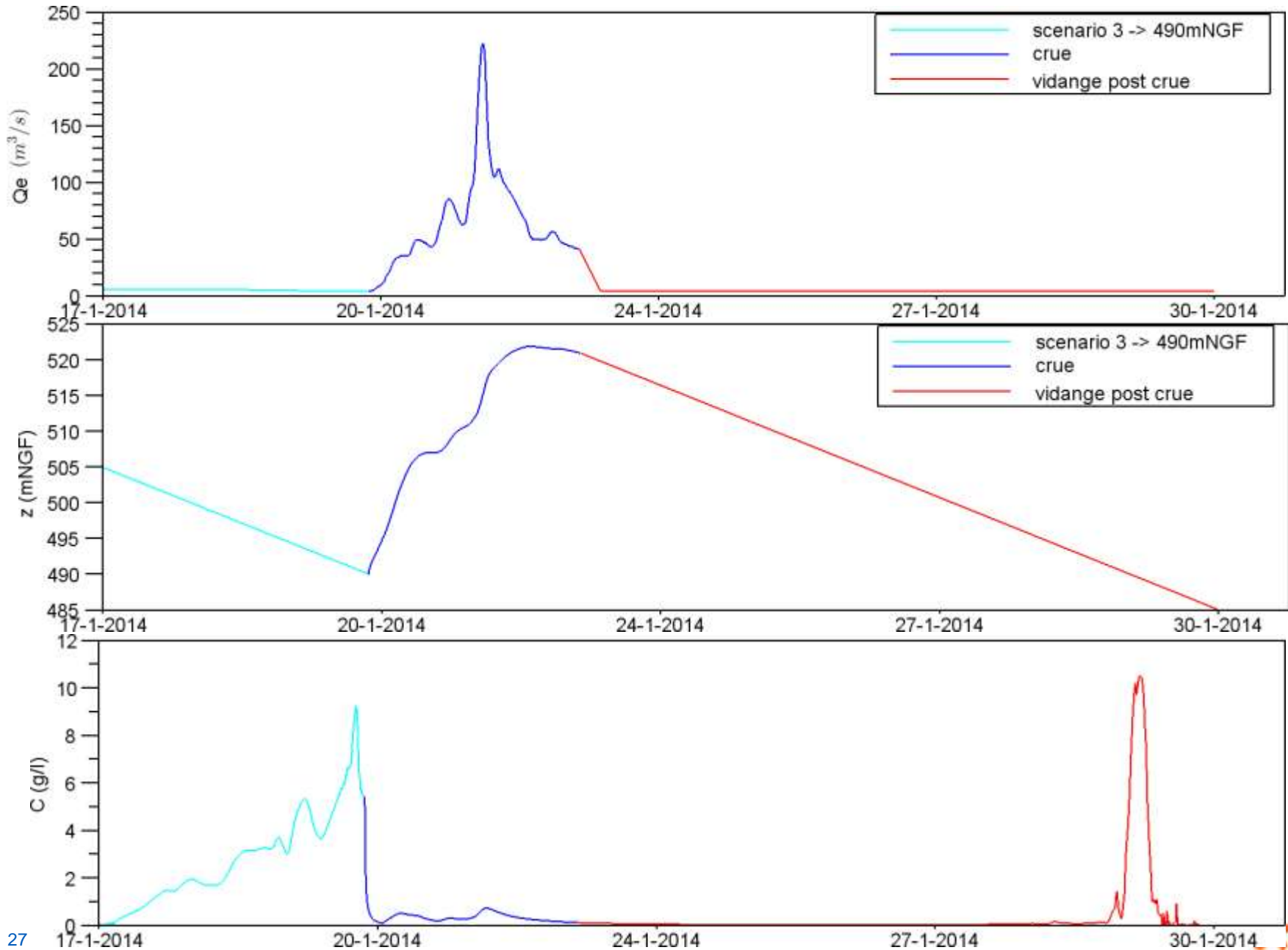
# Emptying scenario: Comparing speed of lowering



# Upstream discharge : no possible in situ control



# What happens if there is a flood ?



# Examples of current research

- ▶ 2011 & 2013 Cemracs Projects

# 2011 Cemracs Project

## ▶ Sediment transport modeling : relaxation schemes for Saint-Venant Exner and three layer models

Emmanuel Audusse, Christophe Chalons, Olivier Delestre, Nicole Goutal, Magali Jodeau, Jacques Sainte-Marie, Jan Giesselmann and Georges Sadaka

EDF, INRIA, UNIV P6

## ▶ SHALLOW WATER AND EXNER EQUATIONS

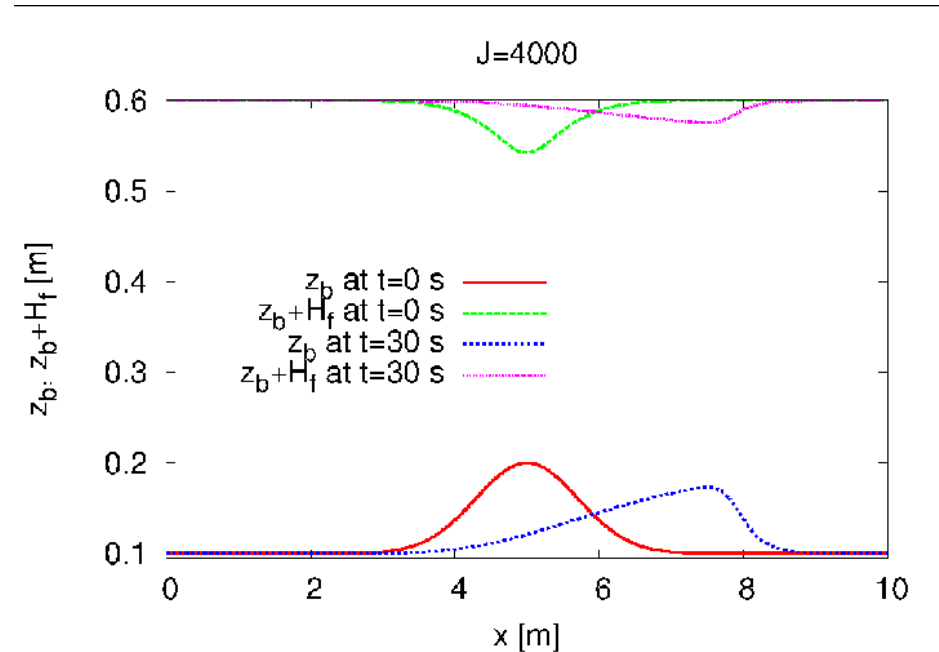
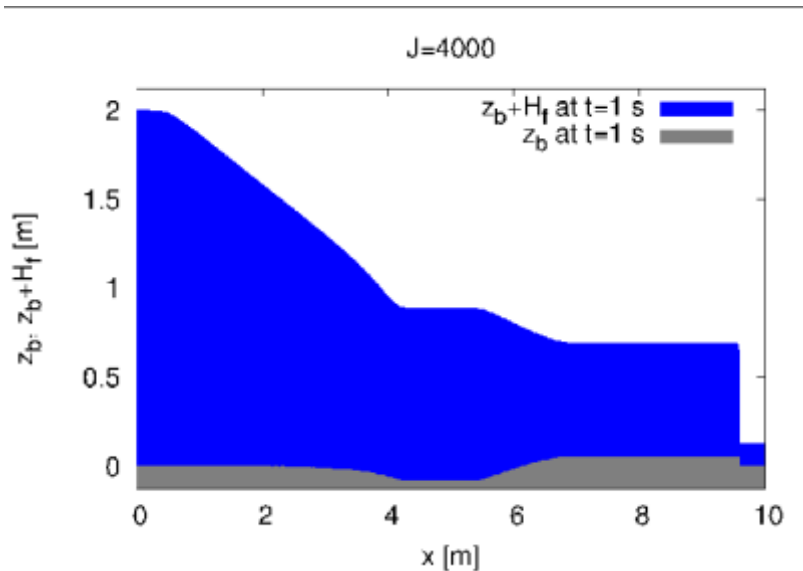
$$\begin{aligned}\frac{\partial H}{\partial t} + \frac{\partial Q}{\partial x} &= 0, \\ \frac{\partial Q}{\partial t} + \frac{\partial}{\partial x} \left( \frac{Q^2}{H} + \frac{g}{2} H^2 \right) &= -gH \frac{\partial z_b}{\partial x}, \\ \rho_s(1-p) \frac{\partial z_b}{\partial t} + \frac{\partial Q_s}{\partial x} &= 0,\end{aligned}$$

## ▶ A RELAXATION SOLVER FOR THE SAINT-VENANT EXNER MODEL

$$\begin{aligned}\frac{\partial H}{\partial t} + \frac{\partial Hu}{\partial x} &= 0 \\ \frac{\partial Hu}{\partial t} + \frac{\partial}{\partial x} (Hu^2 + \Pi) &= -gH \frac{\partial z_b}{\partial x} \\ \frac{\partial \Pi}{\partial t} + u \frac{\partial \Pi}{\partial x} + \frac{a^2}{H} \frac{\partial u}{\partial x} &= \frac{1}{\lambda} \left( \Pi - \frac{gH^2}{2} \right) \\ \frac{\partial z_b}{\partial t} + \frac{\partial \bar{Q}_s}{\partial x} &= 0 \\ \frac{\partial \bar{Q}_s}{\partial t} + \left( \frac{b^2}{H^2} - u^2 \right) \frac{\partial z_b}{\partial x} + 2u \frac{\partial \bar{Q}_s}{\partial x} &= \frac{1}{\lambda} (\bar{Q}_s - Q_s)\end{aligned}$$

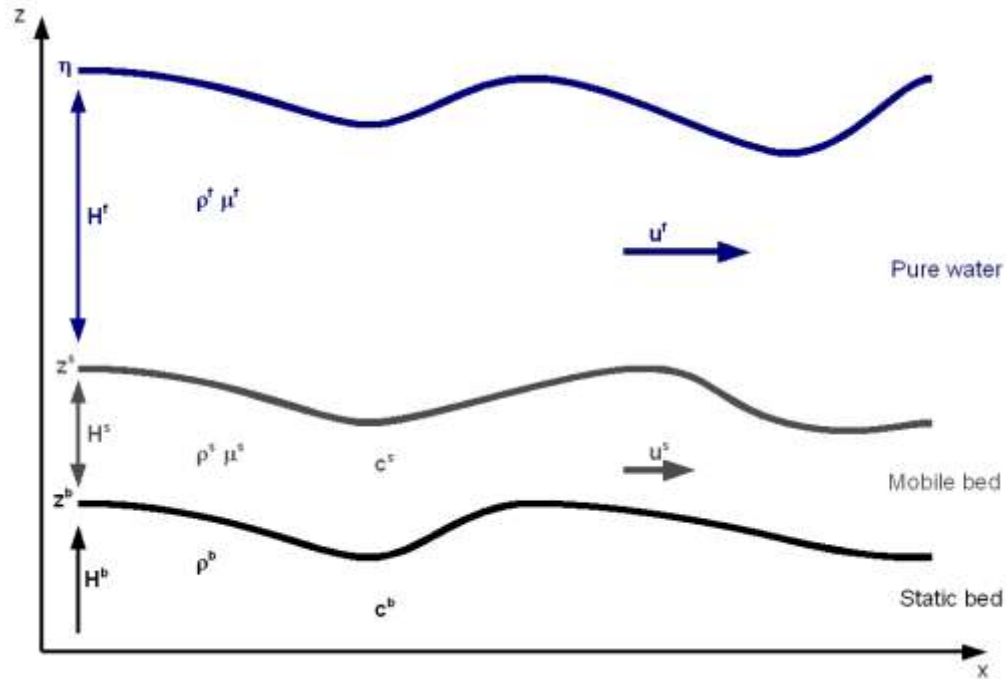
# 2011 Cemracs Project

- ▶ sediment transport modelling : **relaxation schemes for Saint-Venant Exner** and three layer models
- ▶ A RELAXATION SOLVER FOR THE SAINT-VENANT EXNER MODEL : SOME RESULTS



# 2011 Cemracs Project

- ▶ sediment transport modelling : relaxation schemes for Saint-Venant Exner and **three layer models**



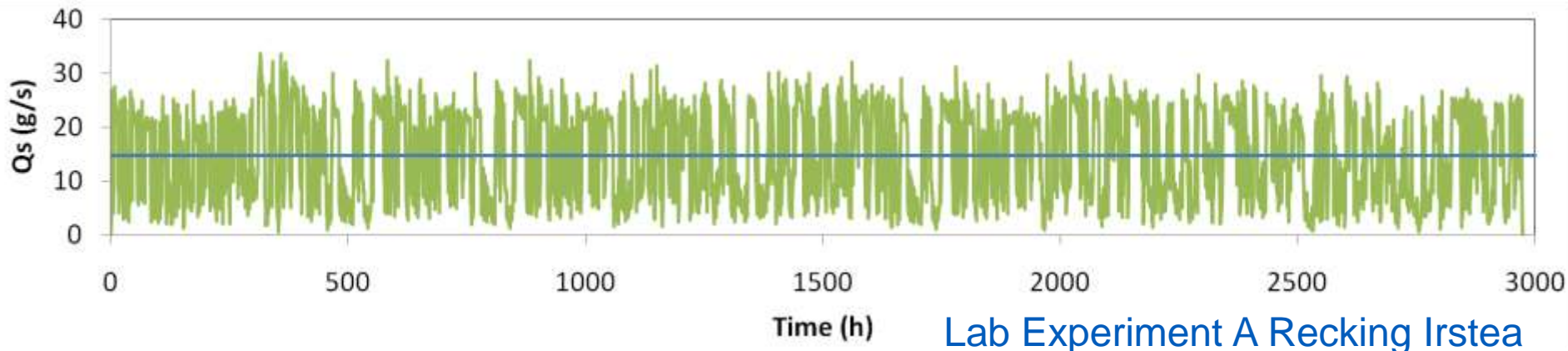
Three Layer Model

# 2013 Cemracs Project

## ► Modeling and simulation of uncertainties in hydraulics and sediment transport

Emmanuel Audusse, Sébastien Boyaval, Yueyan Cao, Nicole Goutal, Magali Jodeau, Philippe Ung  
EDF, UNIV P13, LABORATOIRE ST VENANT

## ► Sediment transport is a stochastic process



## ► How to deal with stochastic properties in numerical modeling ?



# 2013 Cemracs Project

- Modeling and simulation of uncertainties in hydraulics and sediment transport

$$\frac{\partial H}{\partial t} + \frac{\partial Q}{\partial x} = 0,$$

$$\frac{\partial Q}{\partial t} + \frac{\partial}{\partial x} \left( \frac{Q^2}{H} + \frac{g}{2} H^2 \right) = -gH \frac{\partial z_b}{\partial x}, -gHJ$$

$$\rho_s(1-p) \frac{\partial z_b}{\partial t} + \frac{\partial Q_s}{\partial x} = 0,$$

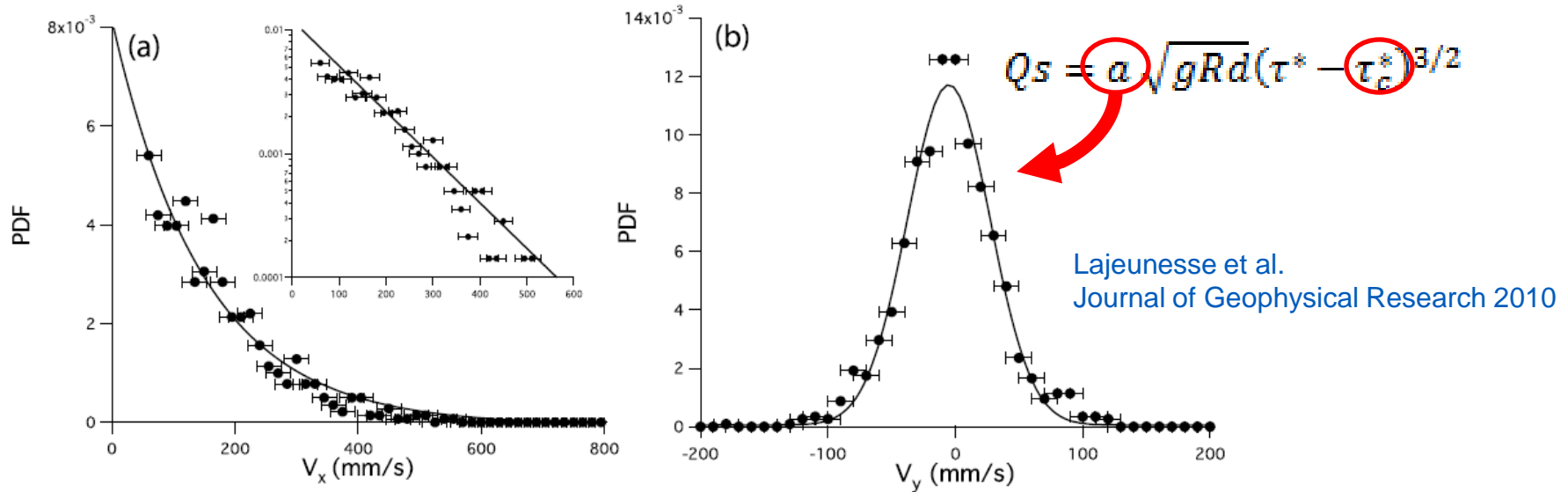
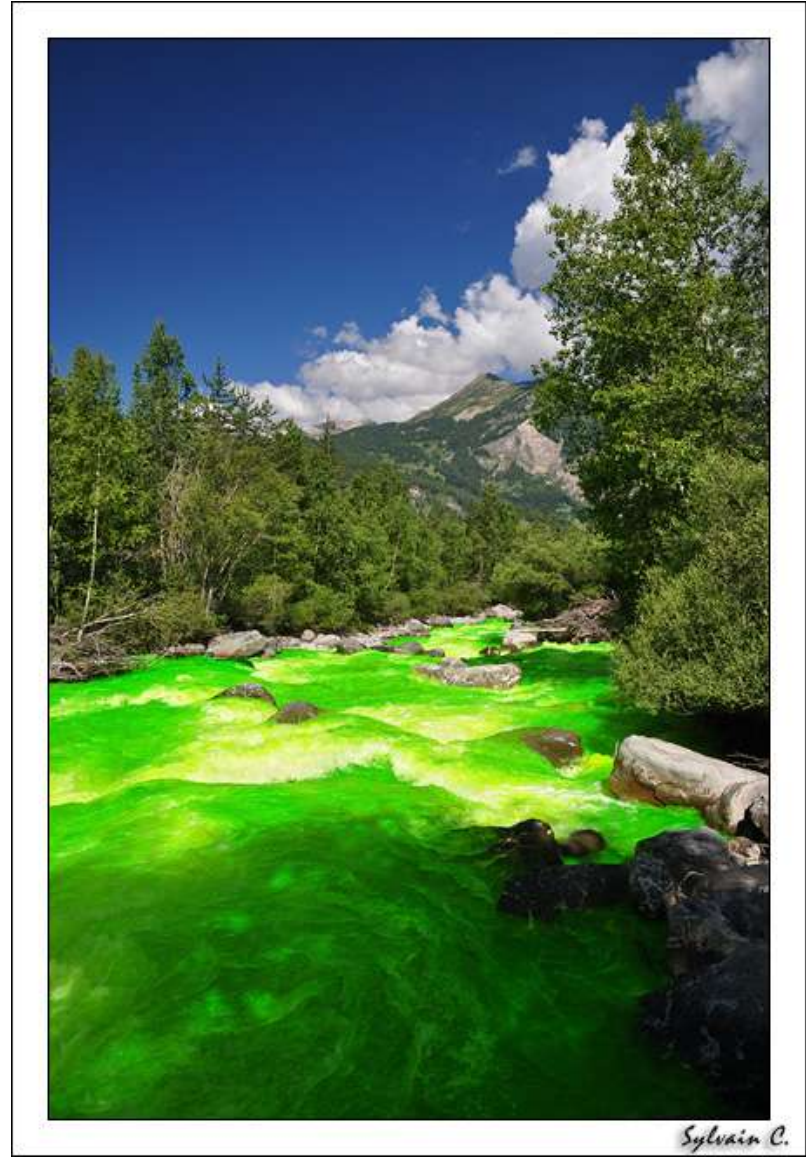


Figure 6. (a) Experimental probability density functions (PDFs) of  $V_x$  for  $\tau^* = 0.103$ ,  $Re_s = 426$ ,  $H/D = 5$ , and  $S = 0.042$ . Inset shows the same data represented on a semilog plot. (b) Corresponding PDFs of  $V_y$ .

Thanks for your attention !



*Sylvain C.*